

# **Pacific Country Report**

## **Sea Level & Climate: *Their Present State***

*Kiribati*

**June 2005**

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**PACIFIC COUNTRY REPORT  
ON  
SEA LEVEL & CLIMATE: THEIR PRESENT STATE**



**KIRIBATI**

**June 2005**

**Executive Summary**

- A SEAFRAME gauge was installed in Betio, Tarawa, Kiribati, in December 1992. It records sea level, air and water temperature, atmospheric pressure, wind speed and direction. It is one of an array designed to monitor changes in sea level and climate in the Pacific.
- This report summarises the findings to date, and places them in a regional and historical context.
- The sea level trend to date is +6.6 mm/year but the magnitude of the trend continues to vary widely from month to month as the data set grows. Accounting for the precise levelling results and inverted barometric pressure effect, the trend is +6.2 mm/year. Nearby gauges, with longer records but less precision and datum control, show trends of -0.68 and +0.27 mm/year.
- Variations in monthly mean sea level include a very small seasonal cycle and were affected by the 1997/1998 El Niño.
- Variations in monthly mean air and water temperature include very small seasonal cycles and were likewise affected by the 1997/1998 El Niño.
- The equatorial location of Tarawa means that it is not subject to tropical cyclones.
- The tsunami caused by the Peru earthquake of June 2001, which registered strongly on many Pacific SEAFRAME gauges, was negligible at Tarawa.

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## **1. Introduction**

As part of the AusAID-sponsored South Pacific Sea Level and Climate Monitoring Project ("Pacific Project") for the FORUM region, in response to concerns raised by its member countries over the potential impacts of an enhanced Greenhouse Effect on climate and sea levels in the South Pacific region, a **SEAFRAME (Sea Level Fine Resolution Acoustic Measuring Equipment)** gauge was installed in Betio, Tarawa, Kiribati, in December, 1992. The gauge has been returning high resolution, good scientific quality data since installation.

SEAFRAME gauges not only measure sea level by two independent means, but also a number of "ancillary" variables - air and water temperatures, wind speed, wind direction and atmospheric pressure. There is an associated programme of levelling to "first order", to determine vertical movement of the sea level sensors due to local land movement. A Continuous Global Positioning System (CGPS) station was installed in Kiribati in August 2002 to determine the vertical movement of the land with respect to the International Terrestrial Reference Frame.

When change in sea level is measured with a tide gauge over a number of years one cannot be sure whether the sea is rising or the land is sinking. Tide gauges measure relative sea level change, i.e., the change in sea level relative to the tide gauge, which is connected to the land. To local people, the relative sea level change is of paramount importance. Vertical movement of the land can have a number of causes, e.g. island uplift, compaction of sediment or withdrawal of ground water. From the standpoint of global change it is imperative to establish absolute sea level change, i.e. sea level referenced to the centre of the Earth, which is to say in the terrestrial reference frame. In order to accomplish this, the rate at which the land moves must be measured separately. This is the reason for the addition of CGPS near the tide gauges.

## **2. Regional Overview**

### **2.1. Regional Climate and Oceanography**

Variations in sea level and atmosphere are inextricably linked. For example, to understand why the sea level at Tuvalu undergoes a much larger annual fluctuation than at Samoa, we must study the seasonal shifts of the trade winds. On the other hand, the climate of the Pacific Island region is entirely ocean-dependent. When the warm waters of the western equatorial Pacific flow east during El Niño, the rainfall, in a sense, goes with them, leaving the islands in the west in drought.

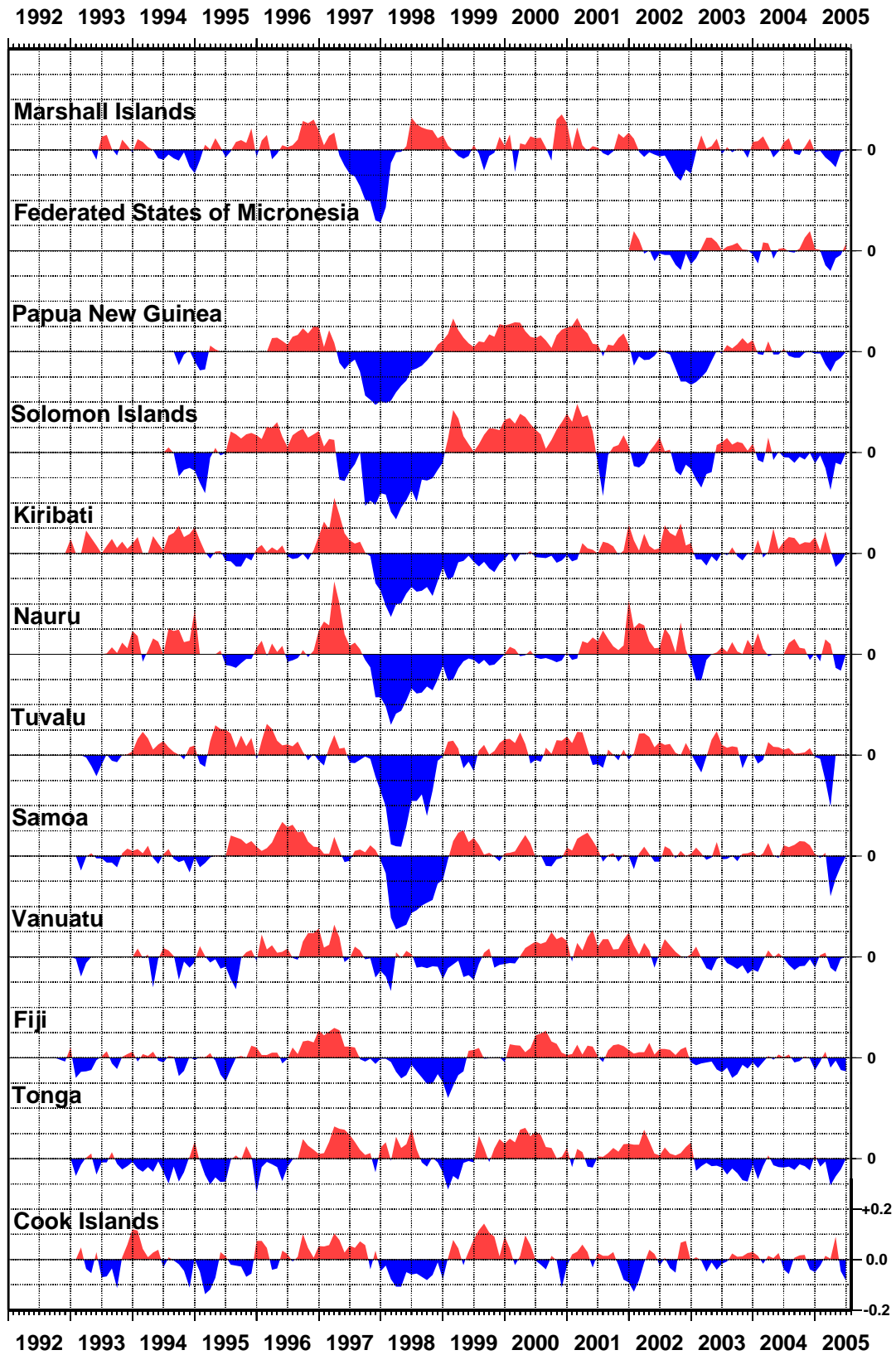
Compared to higher latitudes, air temperatures in the tropics vary little throughout the year. Of the SEAFRAME sites, those furthest from the equator naturally experience the most extreme changes – the Cook Islands (at 21°S) recorded the lowest temperature, 13.1°C, in August 1998. The Cook Islands regularly fall to 16°C while Tonga (also at 21°S) regularly falls to 18°C in winter (July/August).

**Table 1. Range in air temperatures observed at SEAFRAME stations**

<b>SEAFRAME location</b>	<b>Minimum recorded air temperature (°C)</b>	<b>Maximum recorded air temperature (°C)</b>
Cook Islands	13.1	32.0
Tonga	16.0	31.4
Fiji (Lautoka)	16.6	33.4
Vanuatu	16.5	33.3
Samoa	18.7	32.3
Tuvalu	22.8	32.4
Kiribati	22.4	32.9
Nauru	22.4	33.0
Solomon Islands	20.1	34.5
Papua New Guinea	21.5	31.8
Marshall Islands	22.0	31.9
FSM	23.0	31.4

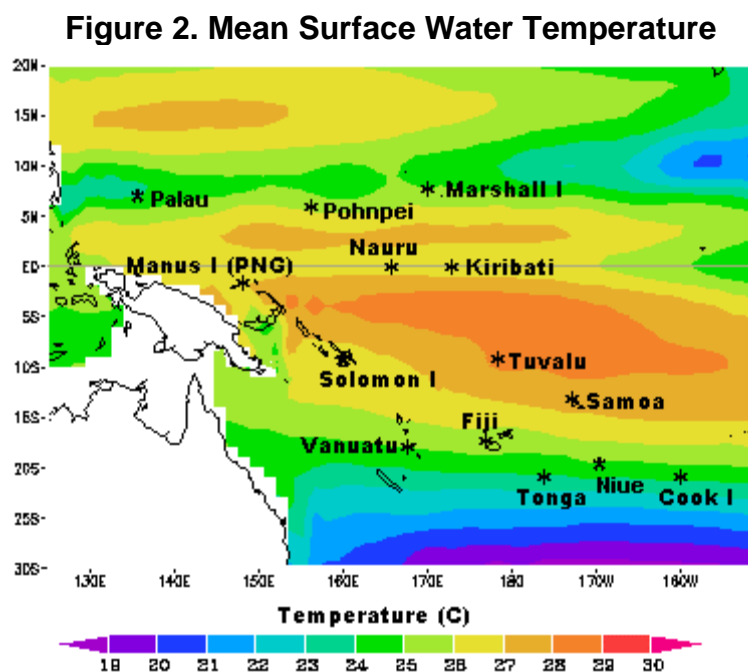
The most striking oceanic and climate fluctuations in the equatorial region are not the seasonal, but interannual changes associated with El Niño. These affect virtually every aspect of the system, including sea level, winds, precipitation, and air and water temperature. Referring to Figure 1, we see that at most SEAFRAME sites, the lowest recorded sea levels appear during the 1997/1998 El Niño. The most dramatic effects were observed at the Marshall Islands, PNG, Nauru, Tuvalu and Kiribati, and along a band extending southeastward from PNG to Samoa. The latter band corresponds to a zone meteorologists call the “South Pacific Convergence Zone” or SPCZ (sometimes called the “Sub-Tropical Convergence Zone”, or STCZ). In Figure 1, we see the effect of the 1997/1998 El Niño on all SEAFRAME stations.

Figure 1. Sea level anomalies\* at SEAFRAME sites



\* Sea level “anomalies” have had tides, seasonal cycles and trend removed from the sea level observations.

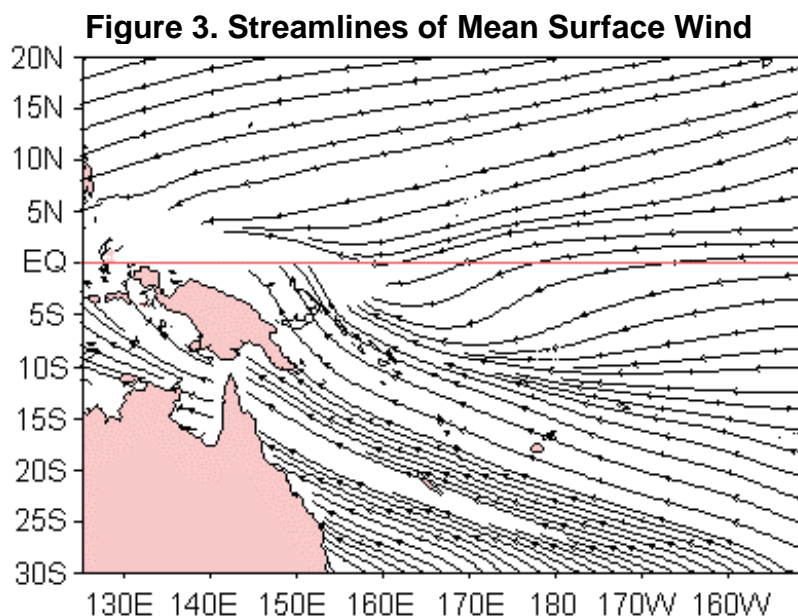
Most Pacific Islanders are very aware that the sea level is controlled by many factors, some periodic (like the tides), some brief but violent (like cyclones), and some prolonged (like El Niño), because of the direct effect the changes have upon their lives. The effects vary widely across the region. Along the Melanesian archipelago, from Manus Island to Vanuatu, tides are predominantly diurnal, or once daily, while elsewhere the tide tends to have two highs and two lows each day. Cyclones, which are fueled by heat stored in the upper ocean, tend to occur in the hottest month. They do not occur within 5° of the equator due to the weakness of the “Coriolis Force”, a rather subtle effect of the earth’s rotation. El Niño’s impact on sea level is mostly felt along the SPCZ, because of changes in the strength and position of the Trade Winds, which have a direct bearing on sea level, and along the equator, due to related changes in ocean currents. Outside these regions, sea levels are influenced by El Niño, but to a far lesser degree.



Note the warm temperatures in the SPCZ and just north of the equator.

The convergence of the Trade Winds along the SPCZ has the effect of deepening the warm upper layer of the ocean, which affects the seasonal sea level. Tuvalu, which is in the heart of the SPCZ, normally experiences higher-than-average sea levels early each year when this effect is at its peak. At Samoa, the convergence is weaker, and the seasonal variation of sea level is far less, despite the fact that the water temperature recorded by the gauge varies in a similar fashion. The interaction of wind, solar heating of the oceanic upper layer, and sea level, is quite complex and frequently leads to unexpected consequences.

The Streamlines of Mean Surface Wind (Figure 3) shows how the region is dominated by easterly trade winds. In the Southern Hemisphere the Trades blow to the northwest and in the Northern Hemisphere they blow to the southwest. The streamlines converge, or crowd together, along the SPCZ.



Much of the Melanesian subregion is also influenced by the Southeast Asian Monsoon. The strength and timing varies considerably, but at Manus Island (PNG), for example, the NW monsoon season (winds from the northwest) runs from November to March, while the SE monsoon brings wind (also known as the Southeast Trade Winds) from May to October. Unlike many monsoon-dominated areas, the rainfall at Manus Island is distributed evenly throughout the year (in normal years).

## **2.2. Historical Sea Level Trends and their Confidence Intervals**

With the great diversity in climatic environments, vertical land movement and ocean variability, one might expect that the relative sea level trends measured at different stations over different time periods may also vary. This is indeed the case and is demonstrated by Table 2, which contains the relative sea level trends from all the 'historical' regional stations. The term 'historical' in this case refers to tide gauges that were installed prior to the South Pacific Sea Level and Climate Monitoring Project. In general, these historical gauges were designed to monitor the sea level variability caused by El Niño and shorter-term oceanic fluctuations rather than long-term sea level change, for which a high level of precision and datum control is required.

**Table 2. Historical Sea Level Data and their Relative Sea Level Trends**

Location	Country	Years of data	Trend (mm/year)	Standard Deviation (mm/year)
Pago Pago	U S Trust	49.7	+1.43	1.5
Rarotonga	Cook Is	22.2	+3.80	3.7
Penrhyn	Cook Is	21.6	+0.89	3.4
Pohnpei	F S of Micronesia	26.9	+0.42	3.7
Kapingamarangi	F S of Micronesia	19.9	-1.04	4.7
Truk	F S of Micronesia	27.6	+1.79	3.3
Guam	U S Trust	50.1	+0.37	1.9
Yap	F S of Micronesia	30.9	-0.20	3.6
Suva	Fiji	24.8	+3.99	3.0
Christmas	Rep of Kiribati	40.3	-0.68	2.2
Kanton	Rep of Kiribati	45.0	+0.26	1.5
Fanning	Rep of Kiribati	16.8	+2.17	5.1
Tarawa	Rep of Kiribati	23.6	-2.24	3.6
Majuro	Rep of Marshall Is	30.8	+2.79	2.6
Enewetok	Rep of Marshall Is	24.5	+1.18	3.3
Kwajalein	Rep of Marshall Is	54.4	+1.13	1.3
Nauru	Rep of Nauru	24.2	-2.03	4.2
Malakal	Rep of Palau	30.1	+0.64	4.0
Honiara	Solomon Is	24.5	-2.21	4.8
Funafuti	Tuvalu	21.6	+0.92	5.1

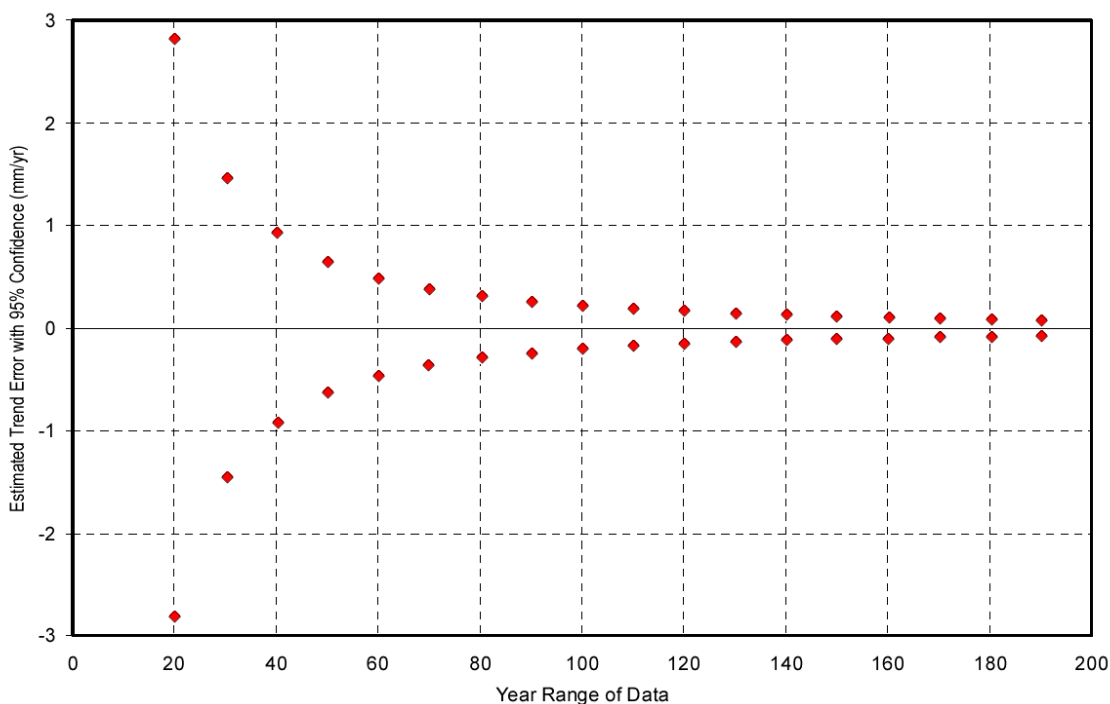
Mean trend: 0.67 mm/year (all data)

Mean trend of data > 25 years: 0.8 mm/year

Data from University of Hawaii as at June 2002

Figure 1 illustrates that sea level can undergo significant short-term fluctuation, yet it can also undergo significant fluctuation on decadal timescales. This sea level variability can affect any estimate of the underlying long-term trend. The expected width of the 95% confidence interval ( $\pm 1.96$  times the standard error) of a linear trend estimate as a function of data length based on the relationship for all National Oceanographic and Atmospheric Administration (NOAA) gauges with a data record of at least 25 years are shown in Figure 4. A confidence interval or precision of 1 mm/year should be obtainable at most stations with 50-60 years of data on average, providing there is no acceleration in sea level change, vertical motion of the tide gauge, or abrupt shifts in trend due to tectonic events.

**Figure 4. 95% Confidence Intervals for Linear Mean Sea Level trends (mm/year) plotted as a function of the year range of data. Based on NOAA tide gauges with at least 25 years of record<sup>1</sup>.**



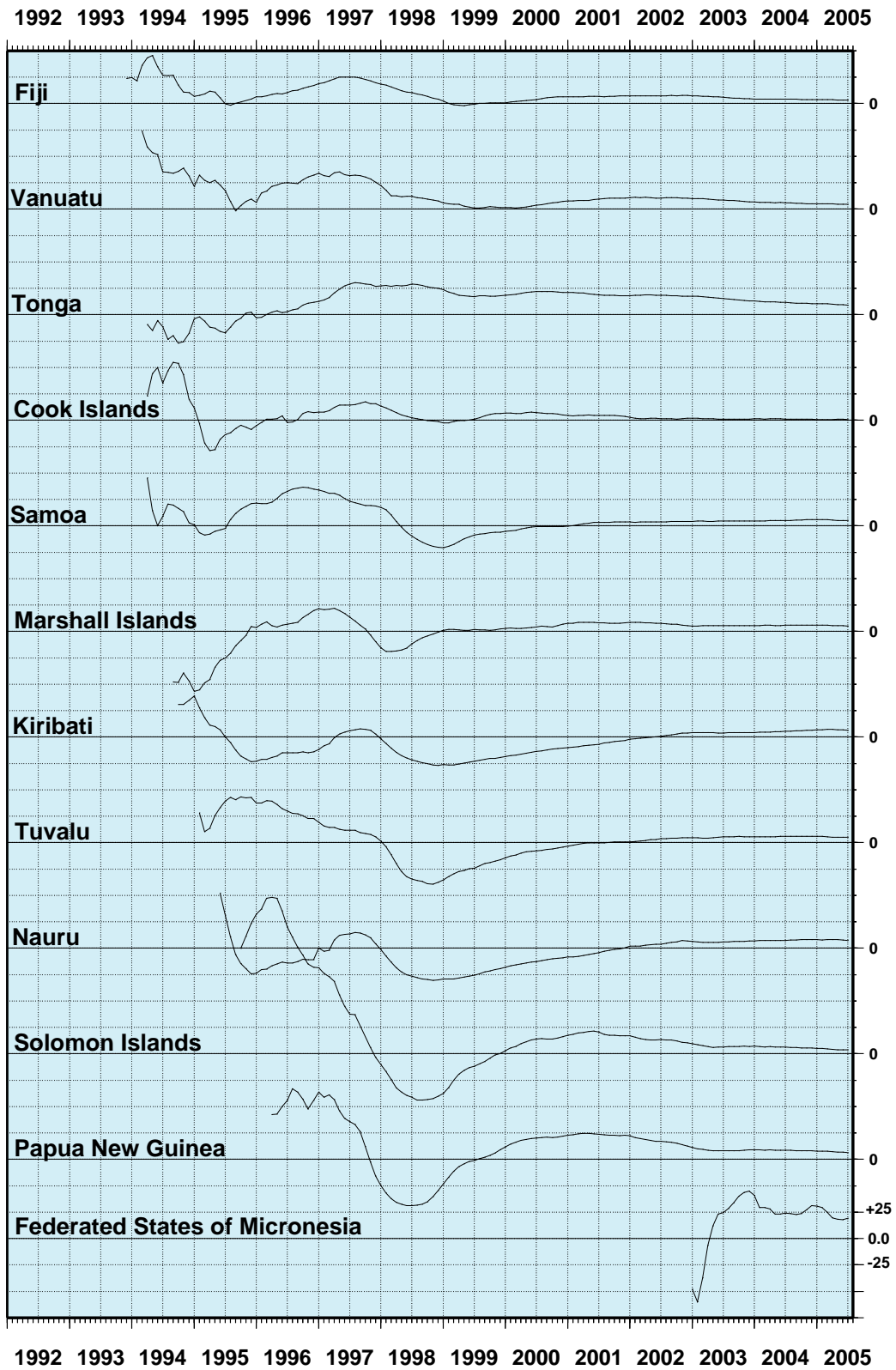
### **2.3. Short-Term Sea Level Trends from SEAFRAME stations**

The importance of precise measurements and vertical datum control for long-term sea level monitoring is integral to the South Pacific Sea Level and Climate Monitoring Project. Nevertheless the data collection program to date has been operating for a relatively short term, and so the trends are still prone to the effects of shorter-term ocean variability (such as El Niño and decadal oscillations). As the data sets increase in length, the trend estimates will begin to reflect longer-term change rather than short-term fluctuations, although Figure 4 shows this transition continues for many decades. Figure 5 shows the evolution of the short-term sea level trends, at SEAFRAME stations, from one year after installation to the present. Please note that the trendlines have not yet stabilised.

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1. Zervas, C. (2001) Sea Level Variations of the United States 1854-1999. NOAA, USA.

Figure 5. Evolution of Short Term Relative Sea Level Trends (mm/year) at SEAFRAME stations.



### 2.3.1 Precise Levelling Summary

Precise datum control is an essential component of the Pacific Project. Precise levelling of the height of the SEAFRAME relative to an array of land-based benchmarks are undertaken periodically, preferably every eighteen months. The precision to which the survey must be performed is dependant on the distance  $K_m$  (km) between the SEAFRAME sensor benchmark and the primary tide gauge benchmark (TGBM) and forms part of Project design specifications.

The precise levelling program essentially monitors the vertical stability of the SEAFRAMES. The rates of vertical movement of the gauges relative to the TGBM (determined by fitting a straight line to the survey results) that is contributing to the observed sea level trend are listed in Table 3. For example, substantial subsidence of the tide gauge at Samoa is occurring at a rate of -1.1 mm/year. Subsidence is also occurring at Marshall Islands and Solomon Islands. The tide gauges at Cook Islands, Fiji, and Vanuatu are rising at 0.3 mm/year with respect to the tide gauge benchmark. The rates of vertical tide gauge movement listed in Table 2.3 are used to correct the observed relative sea level trend estimates in section 2.3.3. **Combined net rate of relative sea level trends.** Additional levelling details for individual countries are provided in section 3.6. **Precise Levelling Results.**

**Table 3. Distance (km), required survey precision (mm), number of surveys and the rate of vertical movement of the SEAFRAME relative to the TGBM.**

Location	$K_m$ (km)	$\pm 2 \sqrt{K_m}$ (mm)	Number of Surveys	Vertical movement (mm/year)
Cook Is	0.491	1.4	8	+0.3
FSM	0.115	0.7	2	N/A
Fiji	0.522	1.4	8	+0.3
Kiribati	0.835	1.8	9	+0.1
Marshall Is	0.327	1.1	8	-0.5
Nauru	0.120	0.7	9	+0.0
PNG	0.474	1.4	7	-0.2
Samoa	0.519	1.4	8	-1.1
Solomon Is	0.394	1.3	4	-0.4
Tonga	0.456	1.4	8	-0.1
Tuvalu	0.592	1.5	8	-0.1
Vanuatu	1.557	2.5	7	+0.3

During Phase 3 of the Project, Continuous GPS stations are being installed on islands in association with the SEAFRAME gauges. Their purpose is to close the final link in establishing vertical datum control – that is, to determine whether the island as a whole (or in some cases the coastal array) is moving vertically. To date, the length of the CGPS data time series is inadequate to make definitive conclusions.

**Table 4. Status of CGPS Installations**

<b>Location</b>	<b>Date of Installation</b>
Samoa	1 July 2001
Cook Is	10 September 2001
Fiji	25 November 2001
Tuvalu	2 December 2001
Tonga	18 February 2002
PNG	1 May 2002
Kiribati	4 August 2002
Vanuatu	11 September 2002
FSM	1 May 2003
Nauru	30 June 2003
Marshall Is	Not yet installed
Palau	Not yet installed
Niue	Not yet installed
Solomon Is	Not yet installed

### 2.3.2. Inverted barometric pressure effect

Another parameter that influences the estimates of relative sea level rise is atmospheric pressure. Known as the inverted barometer effect, if a 1 hPa fall in barometric pressure is sustained over a day or more, a 1 cm rise is produced in the local sea level (within the area beneath the low pressure system). Therefore, if there are trends in the barometric pressure recorded at the tide gauge sites, there will be a contribution to the observed relative sea level trends. The contribution will be a 10 mm/year increase (decrease) in relative sea levels for a 1 hPa/year decrease (increase) in barometric pressure.

Estimates of the contribution to relative sea level trends by the inverted barometric pressure effect at all SEAFRAME sites over the period of the project are listed in Table 5. The estimates are mostly positive, which means relative sea level trends are overestimated without taking the barometric pressure effect into consideration. Corrections to the relative sea level trends are applied section **2.3.3. Combined net rate of relative sea level trends.**

**Table 5. Recent short-term barometric pressure trends expressed as equivalent sea level rise in mm/year based upon SEAFRAME data to June 2005.**

Location	Length of data (months)	Barometric Pressure Contribution to Sea Level Trend (mm/yr)
Cook Is	147	+0.27
FSM	42	-1.24*
Fiji	152	+1.22
Kiribati	145	+0.51
Marshall Is	140	+0.25
Nauru	142	+0.53
PNG	114	+1.54
Samoa	147	+0.41
Solomon Is	127	-0.24
Tonga	148	+0.88
Tuvalu	145	+0.58
Vanuatu	138	+1.41

\*The trend at FSM is from a comparatively short series and therefore varies considerably.

### 2.3.3. Combined net rate of relative sea level trends

The effects of the vertical movement of the tide gauge platform and the inverse barometer effect are removed from the observed rates of relative sea level change and presented in Table 6. These net rates are spatially coherent (with the exception of FSM and Tonga) and consistent with regional sea level trends observed from satellite altimeters over a similar timeframe. The net sea level trend at FSM is comparatively large because it is derived from a comparatively short record. The net sea level trend at Tonga is large in comparison to its neighbouring sites, which could possibly be due to vertical motion of the whole island, but the CGPS record there is still too short (since February 2002) for this motion to be reliably quantified.

**Table 6. The net relative sea level trend estimates after the inverted barometric pressure effect and vertical movements in the observing platform are taken into account.**

Location	Length of data (months)	Sea Level Trend (mm/yr)	Barometric Pressure Contribution (mm/yr)	Vertical Tide Gauge Movement Contribution (mm/yr)	Net Sea Level Trend (mm/yr)
Cook Is	147	+0.7	+0.27	-0.3	+0.7
FSM	42	+19.1 *	-1.24 *	N/A	+20.3*
Fiji	152	+3.2	+1.22	-0.3	+2.3
Kiribati	145	+6.6	+0.51	-0.1	+6.2
Marshall Is	140	+5.3	+0.25	+0.5	+4.6
Nauru	142	+7.5	+0.53	-0.0	+7.0
PNG	114	+6.7	+1.54	+0.2	+5.0
Samoa	147	+5.1	+0.41	+1.1	+3.6
Solomon Is	127	+3.7	-0.24	+0.4	+3.5
Tonga	148	+9.2 **	+0.88	+0.1	+8.2**
Tuvalu	145	+5.0	+0.58	+0.1	+4.3
Vanuatu	138	+4.5	+1.41	-0.3	+3.4

\* The rate at FSM is derived from a comparatively short data record.

\*\*The rate at Tonga is comparatively larger than its neighbours. The vertical motion of the whole island is yet to be quantified.

This section has provided an overview of aspects of the climate and sea level of the South Pacific Sea Level and Climate Monitoring Project region as a whole. The following section provides further details of project findings to date that are relevant to Kiribati.

### **3. Project findings to date - Kiribati**

#### **3.1. Extreme Events**

##### **3.1.1. Tropical Cyclones**

Kiribati, being within 3° of the equator, is not subject to cyclones.

##### **3.1.2. Tsunamis**

A tsunami is a series of waves generated by an impulsive disturbance such as an undersea earthquake, coastal or submarine landslide, volcanic eruption, or asteroid impact. Tsunamis are most commonly generated along tectonic plate margins where earthquakes and volcanoes are found. Due to their association with seismic events tsunamis are also referred to as *seismic sea waves*. The term *tidal wave* is incorrect, as tsunamis have nothing to do with tide generating forces. Tsunami waves may be barely discernible in the open ocean but as they propagate into shallow coastal waters their size may increase significantly.

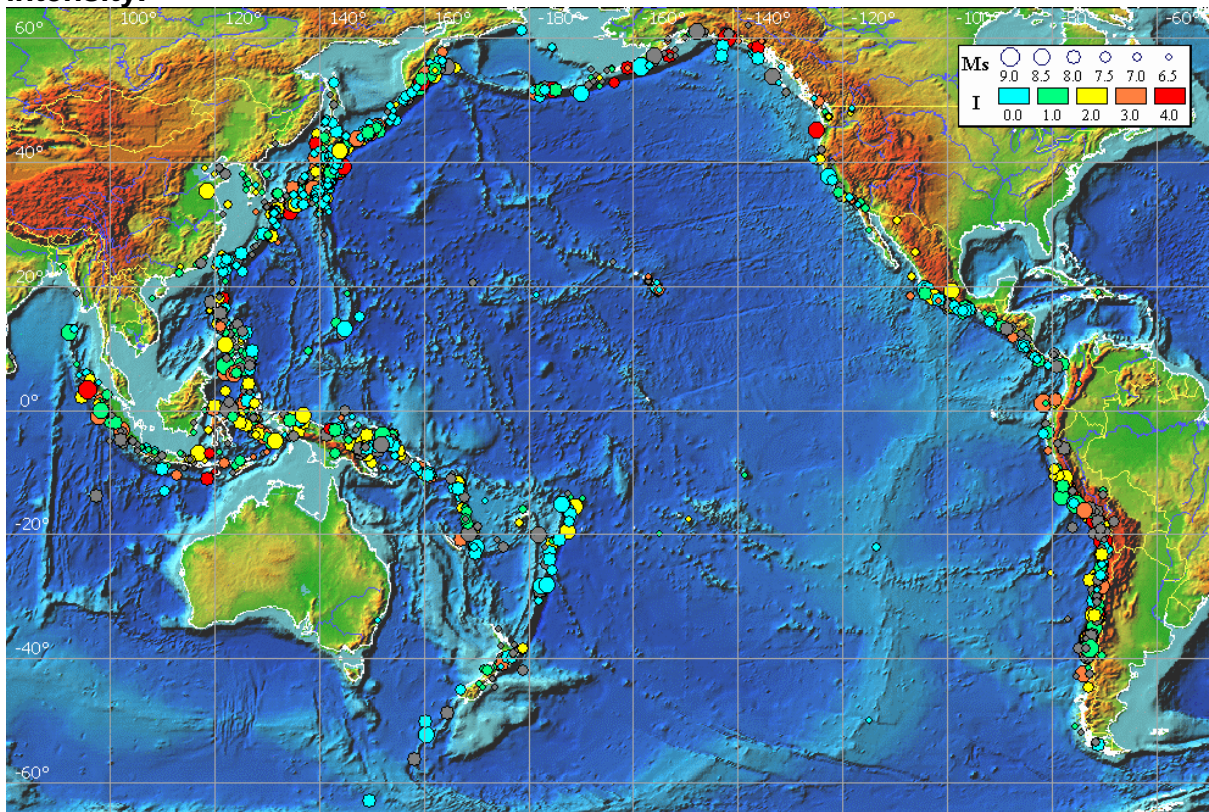
Figure 6 shows the sources of historical tsunami events listed in the *Integrated Tsunami Database for the Pacific and the Eastern Indian Ocean*<sup>1</sup>. A number of tsunamis have been generated in the South Pacific Sea Level and Climate Monitoring Project region. The SEAFRAME tide gauge network has an important role in real time tsunami monitoring and contributes toward the tsunami warning system for the Pacific Ocean.

The historical record reveals that tsunamis have been observed in Kiribati from sources including Russia, Alaska, Chile and Peru. Figure 7 shows the inverse tsunami travel time chart for Tarawa, Kiribati. This chart may be used to provide an estimate of the time taken for a tsunami to arrive at Tarawa from any source location.

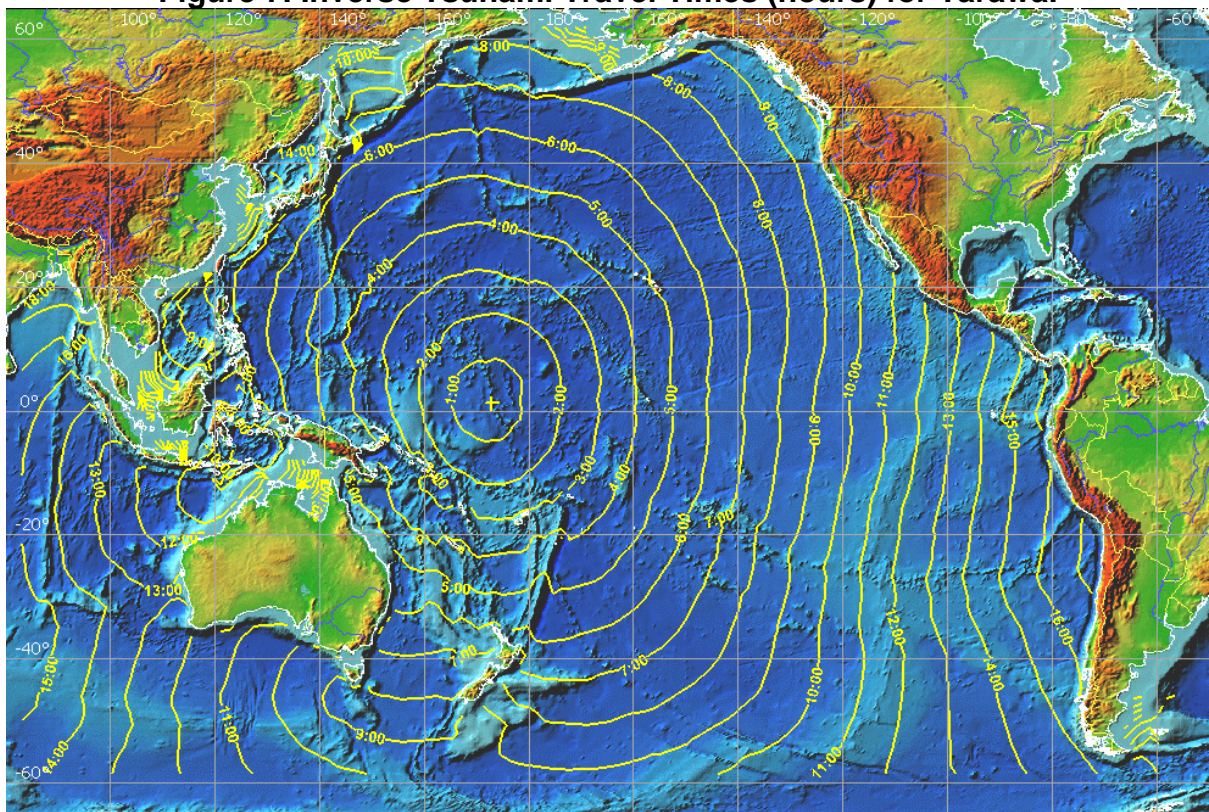
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<sup>1</sup> ITDB/PAC (2004) Integrated Tsunami Database for the Pacific, Version 5.12 of December 31, 2004. CD-ROM, Tsunami Laboratory, ICMG SD RAS, Novosibirsk, Russia.

**Figure 6. Historical Tsunami Events in the Pacific and Eastern Indian Ocean. Circle size indicates earthquake magnitude and colour indicates tsunami intensity.**



**Figure 7. Inverse Tsunami Travel Times (hours) for Tarawa.**



### 3.2. Short-term sea level trend

A fundamental goal of the Project is to establish the rate of sea level change. It has been recognised since the beginning that this would require several decades of continuous, high quality data. However, in response to increasing requests from the region for information regarding the trends as they gradually emerge from the background “noise”, combined with concern that less experienced users might attempt to fit a trend line to the data without properly accounting for processes such as seasonality that can bias the result, the preliminary findings are now being provided. We caution against drawing conclusions prematurely, as illustrated in Figure 5 which shows how the trend develops as more data becomes available.

As at June 2005, based on the short-term sea level trend analyses performed by the National Tidal Centre using the Tarawa SEAFRAME data, a rate of **+6.6 mm per year** has been observed. Accounting for the inverted barometric pressure effect and vertical movements in the observing platform, the sea level trend is **+6.2 mm per year**. By comparison, the Intergovernmental Panel on Climate Change (IPCC) in its Third Assessment Report (IPCC TAR, 2001) estimates that global average long-term sea level rise over the last hundred years was of the order of 1 to 2 mm/yr.

Figure 5 shows how the trend estimate has varied over time. In the early years, the trend appeared to indicate an enormous rate of sea level rise. Later, due to the 1997/1998 El Niño when sea level fell 25 cm below average, the trend actually went negative, and remained so for the next three years. Given the sea level record is still relatively short, it is still too early to deduce a long-term trend.

Figure 8

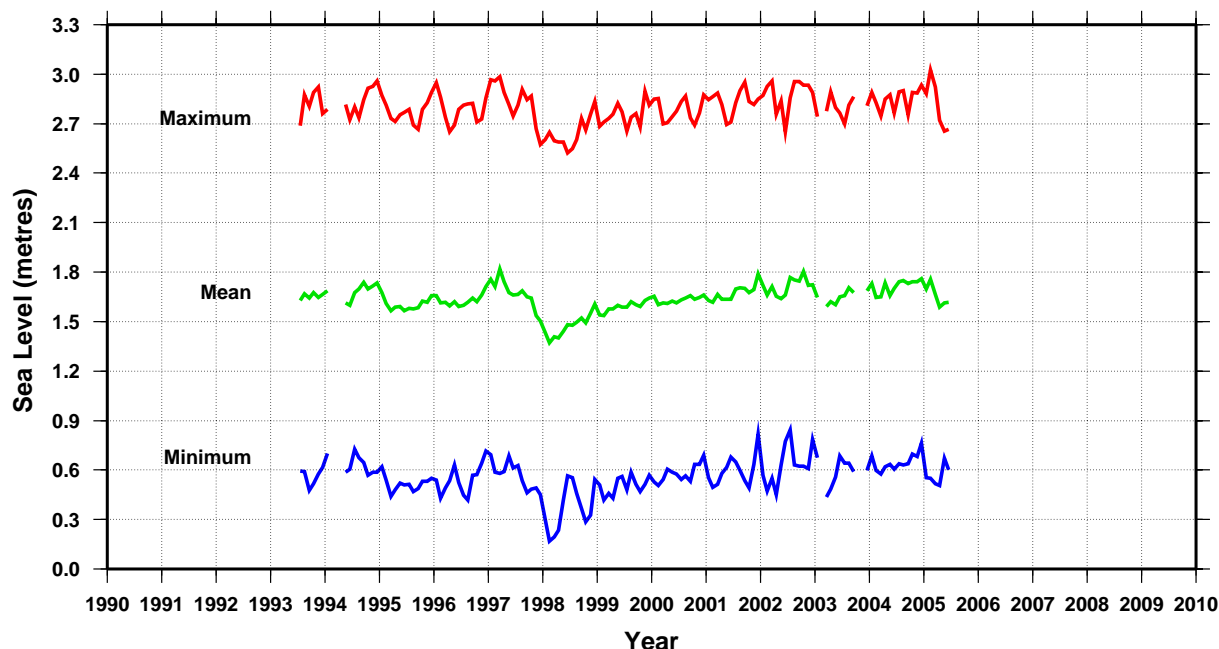


The sea level data recorded since installation is summarised Figure 9. The middle curve (green) represents the monthly mean sea level. The upper and lower curves show the highest and lowest values recorded each month. The monthly mean sea levels remain quite stable throughout the year, although a large fluctuation occurred as a result of the 1997 - 1998 El Niño.

Seasonal cycles are weak. There is normally about a two metre difference between the highest and lowest recorded levels in a given month, a relatively limited range in comparison to many sites. The red curve (highest value for the month) in Figure 9 is the most variable of the three. This relates to the oceanography of the Kiribati region, which is subject to the occasional passage of large-scale equatorial waves. The waves, which are generated by westerly winds on the equator, create large-scale disturbances of the sea surface. These then travel eastwards, temporarily raising the sea surface by several tens of centimetres as they pass. The passage of these waves is most active in the early stages of an El Niño event. The mean sea level over the period was 1.64 metres, with a maximum of 3.02 metres in February 2005, and a minimum of 0.17 metres in February 1998.

**Figure 9**

**Monthly sea level at Betio  
SEAFRAME gauge**



### 3.3. Historical Sea Level Trend Assessment

Longer sea level records are available at Kiribati, from gauges operated by the University of Hawaii (UH) at Tarawa, Kanton, Kiritimati, and Fanning. Records of hourly tide gauge data were kept for about 24, 45, 40, and 17 years respectively. The UH data is summarised in Figures 10 - 13. The respective overall relative trends were **-2.24**, **+0.26**, **-0.68** and **+2.17 mm/year**. These gauges were primarily designed for monitoring El Niño and shorter-term oceanic fluctuations rather than long-term sea level monitoring which requires a high level of precision and datum control. Hence, even with these longer data sets, the trend cannot be established without sizeable uncertainties.

Figure 10

Monthly sea level at Tarawa  
University of Hawaii data

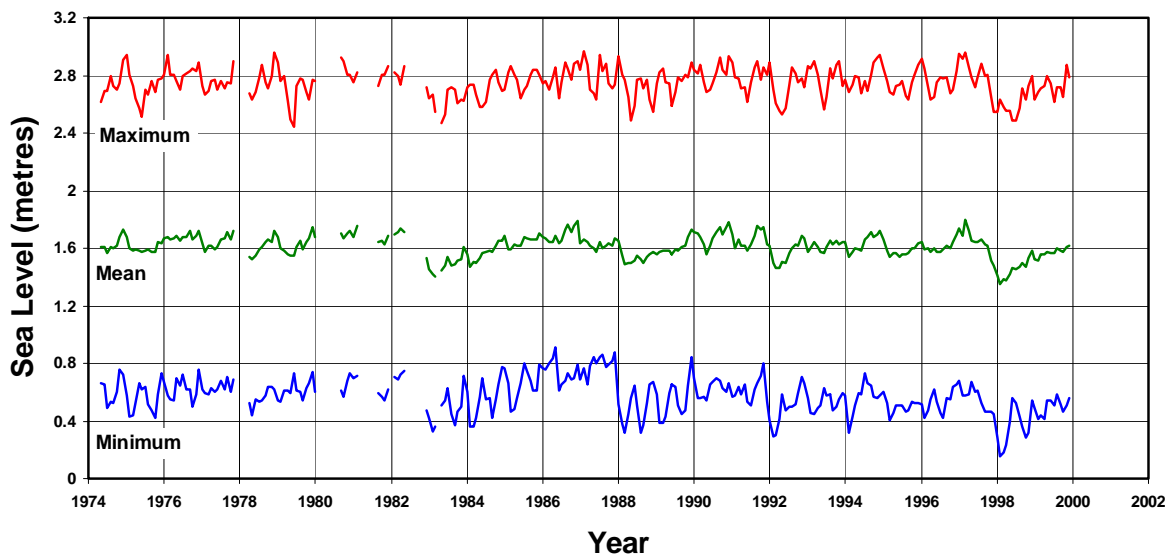


Figure 11

Monthly sea level at Kanton  
University of Hawaii data

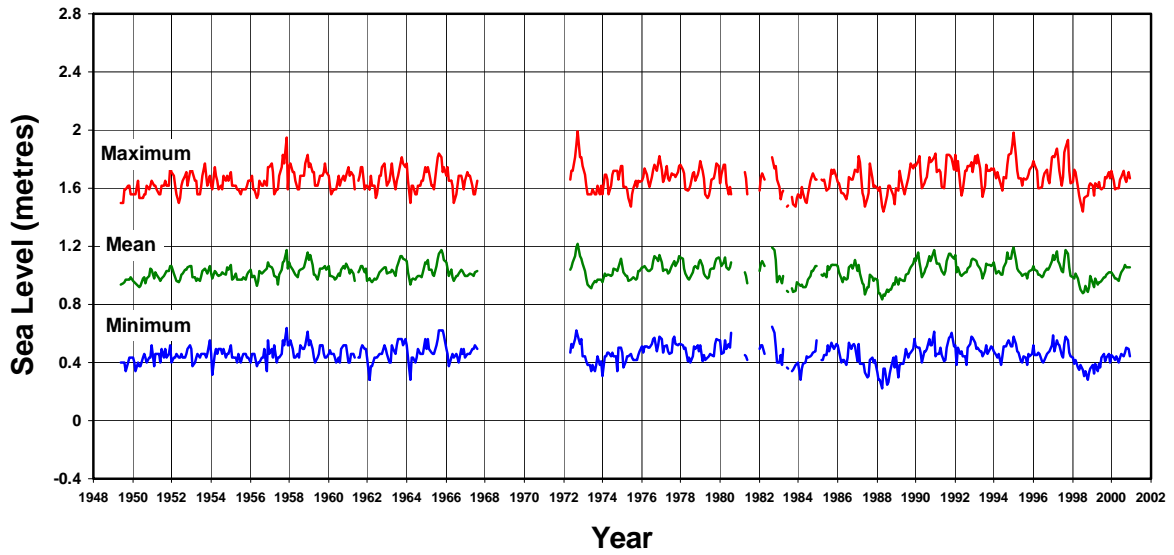


Figure 12

Monthly sea level at Kiritimati Island  
University of Hawaii data

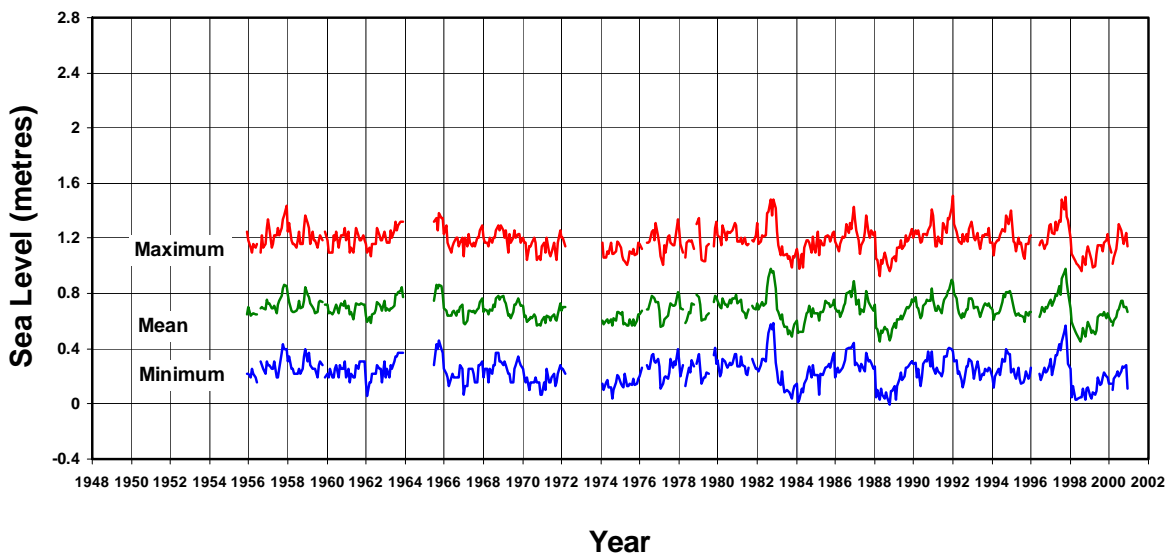
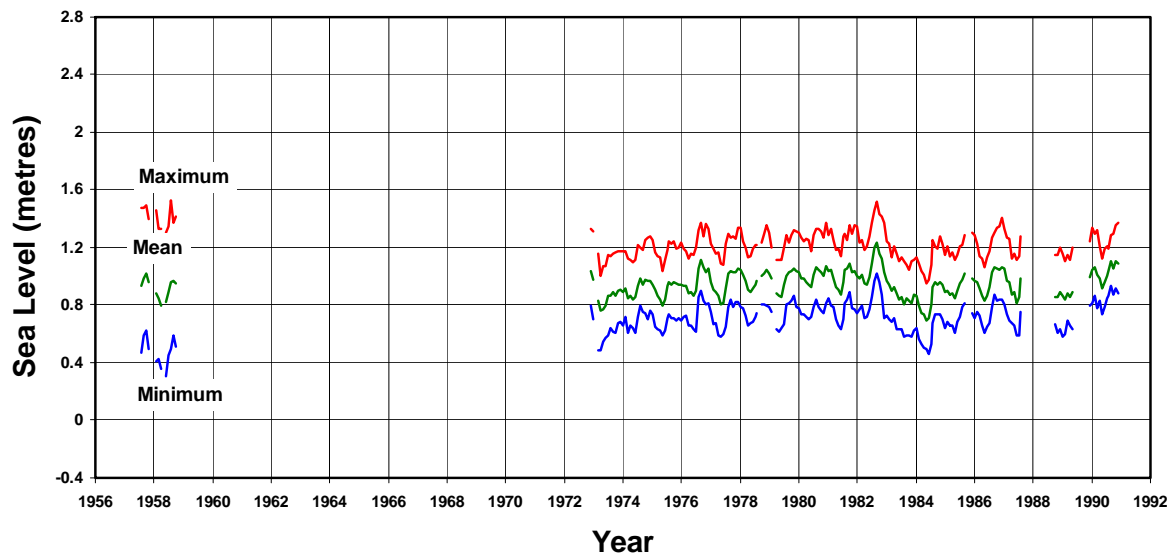


Figure 13

Monthly sea level at Fanning  
University of Hawaii data

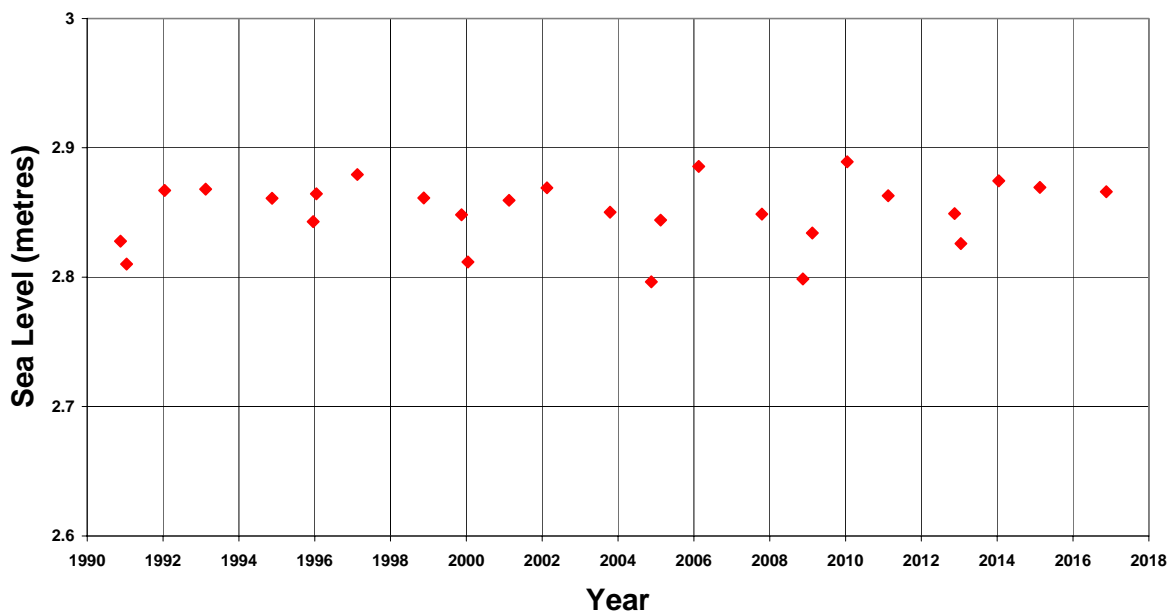


### 3.4. Predicted highest astronomical tide

The component of sea level that is predictable due to the influence of the Sun and the Moon and some seasonal effects allow us to calculate the highest predictable level each year. The highest astronomical tide is the highest sea level that can be predicted under any combination of astronomical conditions, including the proximity of the earth to the sun and the moon. Figure 14 shows that the highest predicted level (2.9 m) over the period 1990 to 2016 will be reached at 17:33 Local Time on the 31 October 2010.

Figure 14

Predicted highest tide each year for Betio



The location of the gauge within the atoll lagoon leads to unique characteristics showing up in the data, such as the effect of solar heating of the lagoon waters. It also shelters the gauge from ocean wave swell, particularly from the east. Swell is caused by surface winds. It is an important source of error in many tide gauges, especially the older conventional gauges with stilling wells.

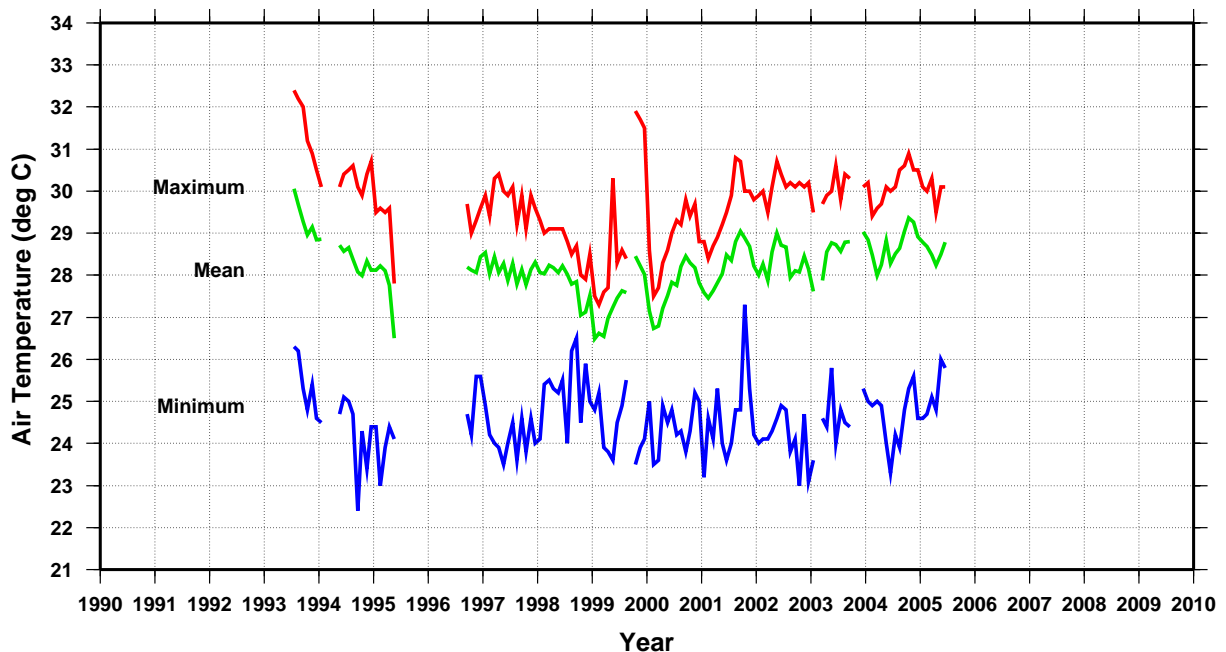
### 3.5. Monthly means of air temperature, water temperature and atmospheric pressure

The data summarised in Figures 15 - 17 follow the same format as the monthly sea level plot: the middle curve (green) represents the monthly mean, and the upper and lower curves show the highest and lowest values recorded each month.

The mean air temperatures at Tarawa have undergone annual fluctuations since the 1997/1998 El Niño, with highest temperatures appearing in September and October. The range is not large, generally having only 4°-6°C separating the highest and lowest temperatures recorded each year. The mean air temperature was 28.2 °C. The highest recorded air temperature was 32.9°C in September 1999, and the minimum was 22.4°C in September 1994. Data was interrupted for over a year, beginning in 1995, due to a combination of power failures at Tarawa and a technical fault with the air thermometer.

Figure 15

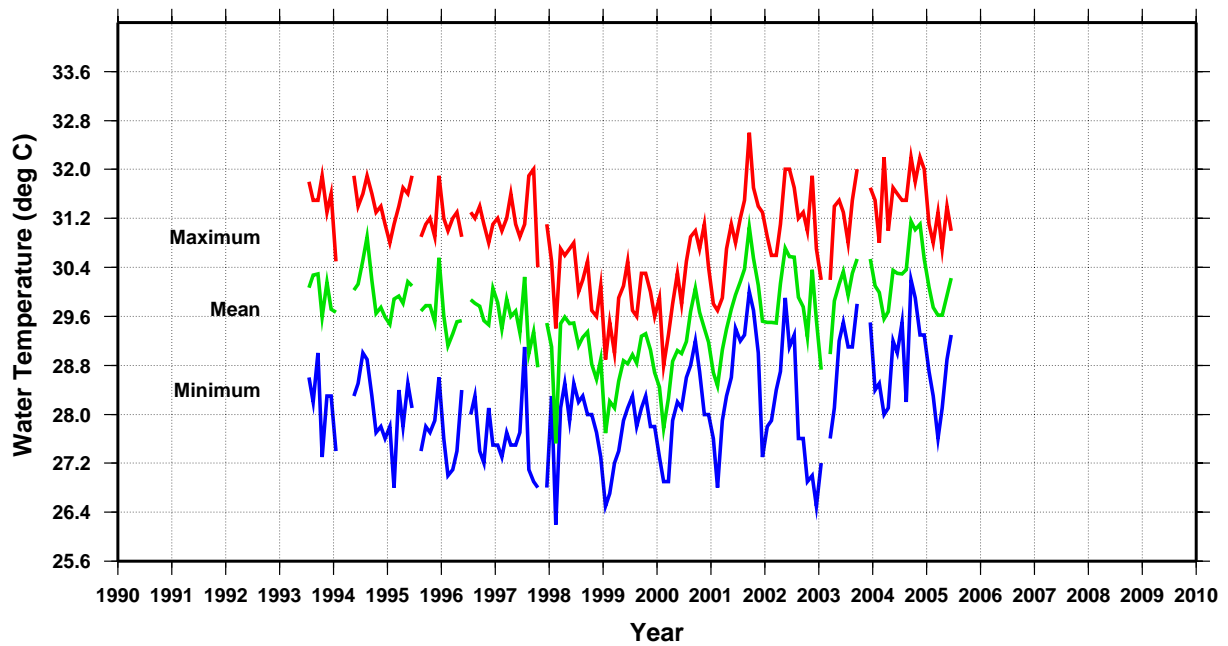
Monthly air temperature at Betio  
SEAFRAME gauge



The annual change in water temperature at Tarawa has about half the range seen in air temperature. The mean water temperature was 29.6°C. The highest recorded water temperature was 32.6°C in September 2001, and the minimum was 26.2°C in February 1998. A notable drop in water temperatures occurred during the 1997/1998 El Niño. Like sea levels, the water temperatures around Kiribati are affected by the passage of large scale, eastward-travelling equatorial waves.

**Figure 16**

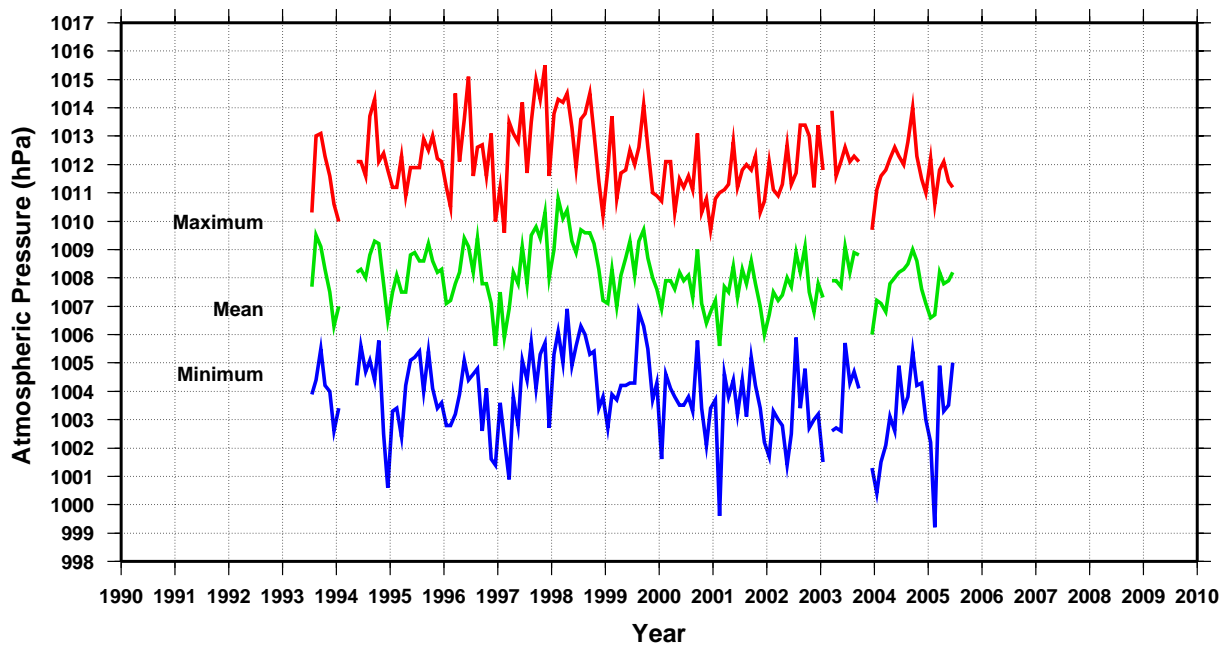
**Monthly water temperature at Betio  
SEAFRAME gauge**



The sea level also responds to changes in barometric pressure. As a rule of thumb, a 1 hPa fall in the barometer, if sustained over a day or more, produces a 1 cm rise in the local sea level (within the area beneath the low pressure system). The monthly atmospheric pressure at Tarawa shows a decline in the years after the El Niño of 1998. The highest pressure recorded was 1015 hPa in November 1997, while the lowest was 999.2 hPa in February 2005.

**Figure 17**

**Monthly atmospheric pressure at Betio  
SEAFRAME gauge**

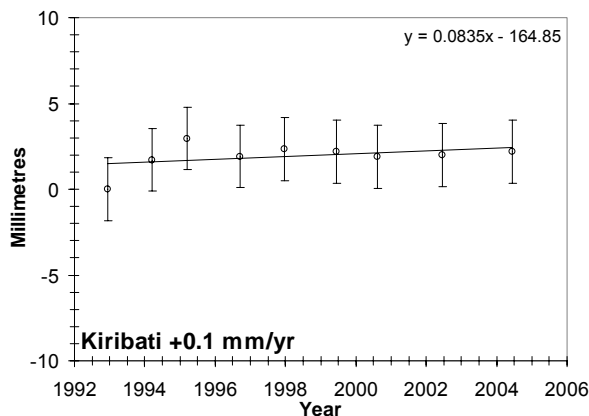


### 3.6. Precise Levelling Results for Kiribati

While the SEAFRAME gauge exhibits a high degree of datum stability, it is essential that the datum stability be checked periodically by precise levelling to an array of deep-seated benchmarks located close to the tide gauge. For example, a wharf normally supports the SEAFRAME, and wharf pilings are often subject to gradual vertical adjustment, which in turn can raise or lower the SEAFRAME.

Precise levelling is carried out on a regular 18-monthly cycle between the SEAFRAME Sensor Benchmark and an array of at least six deep benchmarks. The nearest stable benchmark is designated the “Tide Gauge Benchmark (TGBM)”, and the others are considered the “coastal array”.

Figure 18 summarises the most important survey information being the movement of the SEAFRAME Sensor benchmark relative to the TGBM. The graph does not include the results for the other benchmarks on the coastal array. In this graph, each survey is plotted relative to the first. Thus, the second survey at Kiribati found that the SEAFRAME Sensor benchmark had *risen* relative to the TGBM by 2 mm, and it has continued to rise at an average rate of 0.1 mm/year. This movement is accounted for in section **2.3.3. Combined net rate of relative sea level trends.**



**Figure 18. Movement of SEAFRAME Sensor relative to the Tide Gauge Bench Mark**



**Levelling of SEAFRAME Sensor benchmark. Photo credit: Steve Turner, NTC.**

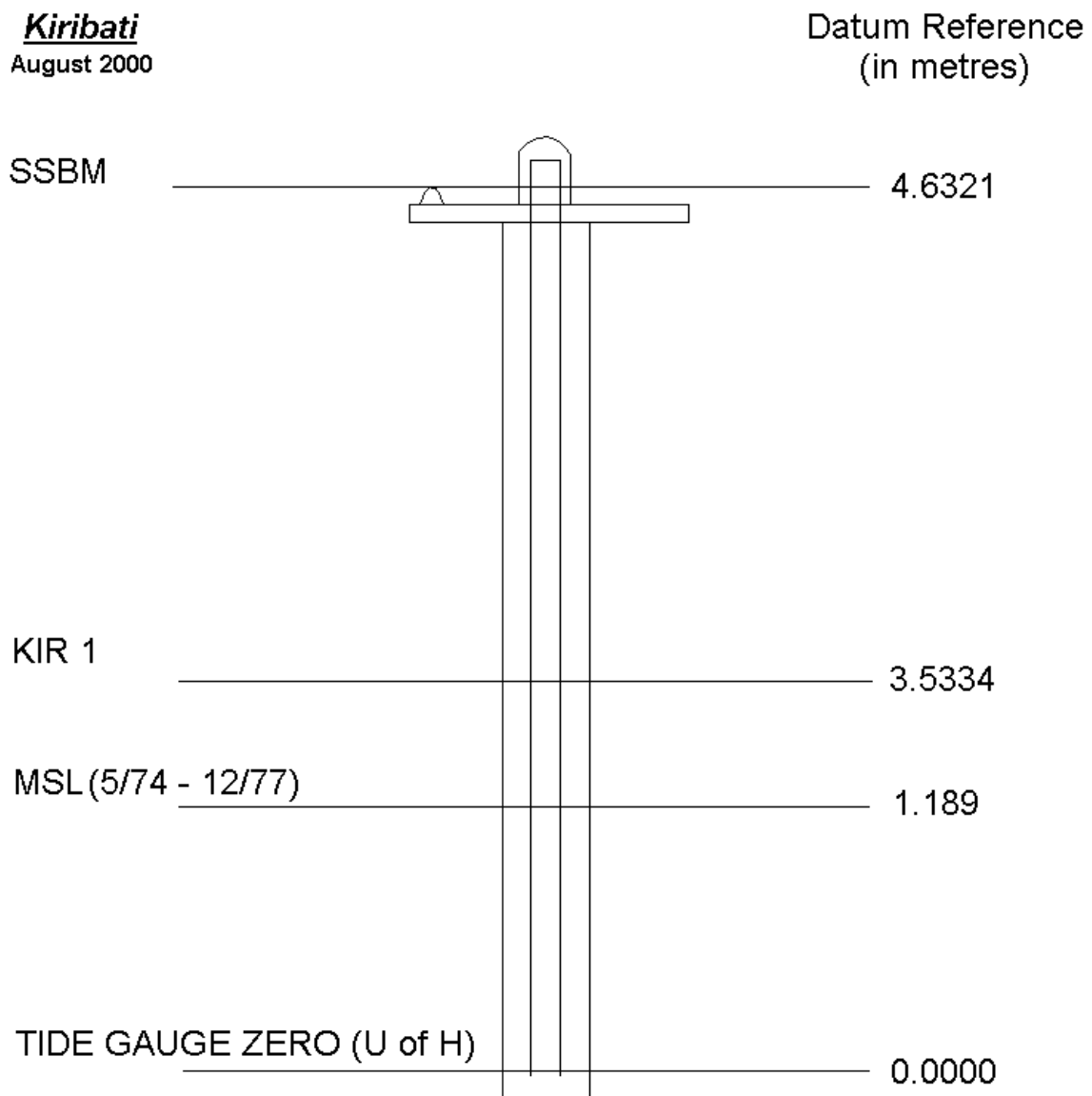
## Appendix

### **A.1. Definition of Datum and other Geodetic Levels at Tarawa**

Newcomers to the study of sea level are confronted by bewildering references to “Chart Datum”, “Tide Staff Zero”, and other specialised terms. Frequently asked questions are, “how do NTC sea levels relate to the depths on the marine chart?” and “how do the UH sea levels relate to NTC’s?”.

Regular surveys to a set of coastal benchmarks are essential. If a SEAFRAME gauge or the wharf to which it is fixed were to be damaged and needed replacement, the survey history would enable the data record to be “spliced across” the gap, thereby preserving the entire invaluable record from start to finish.

**Figure 19**



The word “datum” in relation to tide gauges and nautical charts means a reference level. Similarly, when you measure the height of a child, your datum is the floor on which the child stands.

“Sea levels” in the NTC data are normally reported relative to “Chart Datum” (CD), thus enabling users to relate the NTC data directly to depth soundings shown on marine charts – if the NTC sea level is +1.5 metres, an additional 1.5 metres of water may be added to the chart depths. Unfortunately, at Tarawa the original benchmark used for the marine surveys is unrecoverable, so it is not possible to place CD on Figure 19. In the absence of a known CD, NTC has chosen to refer sea level to the older UH datum, or “Tide Staff Zero”. With this choice, the Mean Sea Level of either data set is close (though not necessarily identical).

Mean Sea Level (MSL) in Figure 19 is the average recorded level at the gauge over the three and a half year period 1974/1977 (as indicated). The 1974/1977 MSL at Tarawa was 1.189 metres above the UH Tide Staff Zero (and the SEAFRAME zero level).

